

Article

Dynamics of soil CO₂ efflux in three tropical dry deciduous forests of Central Indian landscape

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Abstract

Dynamics of soil CO₂ efflux in tropical dry deciduous forests is imperative to know their contribution in regulating the regional and global carbon (C) cycles. In this study, three forest types: dry deciduous teak (DDTF), dry deciduous mixed (DDMF) and *Boswellia* (BF) forests were selected to measure the dynamics of soil CO₂ efflux and its driving factors. Significantly ($p < 0.001$) higher mean monthly CO₂ efflux was recorded in DDMF (626.1 ± 9.1 mg CO₂/m²/h) while it was lowest in BF (122.3 ± 5.0 mg CO₂/m²/h) and DDTF (142.8 ± 6.9 mg CO₂/m²/h) forest types, respectively. The CO₂ efflux peaked during the rainy season (mean 551.1 ± 63.5 mg CO₂/m²/h, DDMF) followed by summer (363.7 ± 68.6 mg CO₂/m²/h, DDTF) and the lowest in winter (181.8 ± 36.3 mg CO₂/m²/h, BF) season. Significantly ($p < 0.05$) lower soil temperature (T_s) and higher soil moisture (M_s) content were observed in BF and DDMF forest types, respectively. The cumulative annual soil CO₂ efflux was highest in DDMF (4625.2 mg CO₂/m²/yr) and lowest in BF (3536.7 mg CO₂/m²/yr). Soil CO₂ efflux was significantly positively correlated with T_s ($R^2 = 0.49$), M_s ($R^2 = 0.59$). This study will provide an understanding of the dynamics of soil CO₂ efflux among tropical dry deciduous forest types in the Central Indian landscape and identify the roles of different drivers in soil CO₂ efflux.

Keywords soil CO₂ efflux; seasonal dynamics; soil moisture; tropical dry deciduous forests; Central India.

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1 Introduction

The carbon (C) cycle is a biogeochemical process that exchanges fluxes of C between different reservoirs, such as the atmosphere, geosphere, hydrosphere and pedosphere (Riebeek, 2011). Soil respiration plays a significant role in global C cycle, releasing 98 Pg C per year into the atmosphere (Bilandžija et al., 2016; Zhao et al., 2017). Globally, carbon dioxide (CO₂) efflux is estimated to be 98 ± 12 Pg C/yr or 85 Pg C/yr if agricultural areas are excluded and is increasing at a rate of 0.1 Pg C/yr (Bond-Lamberty and Thomson, 2010). It is considered as the second largest C flux between the terrestrial ecosystems and the atmosphere (Hanson et al., 2000). Soil CO₂ flux consists of two major C fluxes, i.e., autotrophic respiration of plant roots and heterotrophic respiration through the soil microbial activities and is affected by a multitude of environmental factors (Chen et al., 2017; Zhou et al., 2013). It is estimated that about 90% of the soil CO₂ emissions is performed by the soil microflora via decomposition of organic matter (Reichle et al., 1975). Root respiration can also contribute to about 50% of the total respiration in soil (Pregitzer et al., 2007). It is therefore important to understand the dynamics of soil respiration in any ecosystem to mitigate climate change.

Tropical forests play a significant role in global C cycling as it comprises about 40% of all global terrestrial biomass C stocks and adds up to 67% of the overall soil CO₂ flux (Field et al., 1998; Phillips et al., 1998; Bond-Lamberty and Thomson, 2010b). Thereby, estimations of soil respiration from tropical forests are very useful in understanding total metabolic activities that occur in the soil, concentrations of carbon fluxes via soil, and interrelations among soil and air. Howard and Howard (1993) have examined the relationships that exist among soil CO₂ evolution, moisture content and temperature for various soil types. Various abiotic and biotic factors such as soil temperature, soil moisture (Arroyo and Wood, 2021; Bao et al., 2016), soil pH (Andersson and Nilsson, 2011), availability of C substrates for microorganisms (Okello et al., 2023), soil microbial activity (Tang et al., 2018), soil fertility (Butnor et al. 2003), plant photosynthetic activity (Zhang et al., 2013a, 2013b), and soil organisms (Rai and Srivastava 1981) affect the rate of soil CO₂ efflux.

Even small changes in soil CO₂ fluxes across vast areas can have a considerable impact on atmospheric CO₂ concentration and provide potential positive feedback to global warming (Schlesinger and Andrews, 2000). Comprehensive data on soil CO₂ efflux and its control variables are therefore crucial to frame the ecosystem C budget and identify the response of soil to global warming and climate change in different ecosystems (Buchmann, 2000; Han et al., 2007). With this backdrop, the present study aimed: i) to assess the seasonal and monthly variations in soil CO₂ efflux, ii) to understand the impacts of soil temperature and soil moisture on soil CO₂ efflux, and iii) to analyze the relationships between CO₂ efflux with soil temperature (Ts) and soil moisture (Ms) in three different tropical dry deciduous forest types of Sagar, Madhya Pradesh.

2 Study area and Methodology

2.1 Study area

The present study was undertaken in three tropical dry deciduous forest types viz., dry deciduous teak forest (DDTF, Site I), dry deciduous mixed forest (DDMF, Site II) and *Boswellia* forest (BF, Site III) in Sagar, Madhya Pradesh (Table 1, Fig. 1) which are a part of lower Vindhyan range of Central India, and is situated at an average height of 420 m a.s.l. The area has a hot dry summer from April to June, followed by a monsoon season from July to September and a cool and relatively dry winter from October to March. The area receives an annual average rainfall of 1187.6 mm of which the rainy months contribute approximately 90%. The mean annual minimum and maximum temperatures vary between 10°C (January) and 42.7°C (May) respectively (WorldClim, 2020). As per Champion and Seth's Classification (Champion and Seth, 1968), the forest in the area belongs to group 4b. Vegetation and soil CO₂ efflux were done in three forest ranges of Sagar district, viz. Shahgarh, Nanakpur and Heerapur by laying 14 plots of 60 m × 20 m in each study site/forest type following

standard methods (Misra, 1968). In each plot, the diameter of all the tree individuals (≥ 10 cm diameter at breast height of 1.37 m (DBH)) were recorded. The predominant tree species in the study sites of the three different forest types are *Tectona grandis*, *Butea monosperma* and *Lagerstroemia parviflora* in DDTF, *Tectona grandis*, *Terminalia tomentosa* and *Lagerstroemia parviflora* in DDMF, and *Boswellia serrata*, *Tectona grandis* and *Lagerstroemia parviflora* in BF. A total of 25 tree species (24 genera, 14 families) were recorded from the above-said three forest types. Diversity indices (Shannon's diversity and Simpson's dominance) were calculated using PAleontological STatistics software (PAST; Hammer et al., 2001; Natural History Museum, University of Oslo).

Table 1 Study site characteristics of the three different forest types (DDTF - dry deciduous teak forest, DDMF - dry deciduous mixed forest and BF - *Boswellia* forest) of Sagar, Madhya Pradesh.

Parameter	Forest type		
	DDTF	DDMF	BF
Latitude (°)	24.212 - 24.226	24.130 - 24.241	24.240 - 24.243
Longitude (°)	79.111 - 79.113	79.144 - 79.145	79.121 - 79.133
Altitude (m a.s.l.)	401.6	371.8	486.2
Slope (°)	0.26 - 1.25	0.27 - 0.84	0.40 - 1.06
Aspect (°)	262.2 - 354.3	105.3 - 187.3	213.8 - 245.7
No. of plots	14	14	14
Species richness	14	23	15
Genera	14	22	15
Families	9	13	10
Shannon's index (H)	1.14	2.08	1.76
Dominance index (D)	0.46	0.19	0.27
Evenness index	0.22	0.34	0.39
Total tree density	8192	12032	7272
Density (No./ha)	585.1 \pm 35.7	859.4 \pm 93.1	519.4 \pm 35.9
Total tree basal area	287.9	287.1	413.3
Basal area (m ² /ha)	20.6 \pm 1.7	20.5 \pm 1.4	29.5 \pm 1.6
Maximum tree DBH (cm)	60.8	54.1	62.9
Mean tree DBH (cm)	18.9	15.7	24.3

DBH: Diameter at breast height.

2.2 Estimation of soil respiration

Soil CO₂ efflux was measured by alkali absorption method (Gupta and Singh, 1977), at three different forest types, using plastic jars, inserted 10 cm into the ground. Three replicate sub-plots in each forest type were

selected for the measurement of soil CO₂ efflux. Five replicates of the plastic jars were set up in each sub-plot, and one set of three control plastic jars with airtight lids in each sub-plot. Before each plastic jar was fixed, the vegetation falling within the plastic jar was clipped at the base with the help of scissor. A 50 ml beaker containing 20 ml 0.5 N NaOH was placed in a thin wire tripod stand that holds the jar off the ground by about 2 cm. The alkali was titrated against 1N HCL after 24 hours of absorption period to avoid diurnal variations (Joshi et al., 1991; Harris and van Bavel, 1957). The jars were placed randomly, and on each sampling date the soil moisture was measured by gravimetric method up to 10 cm soil depth. The CO₂ evolved during the experiment was calculated by following the formula of Joshi et al. (1991).

$$\text{mg CO}_2 = V \times N \times 22$$

where V represents titration of the blank minus the sample titration and N is the normal acid value.

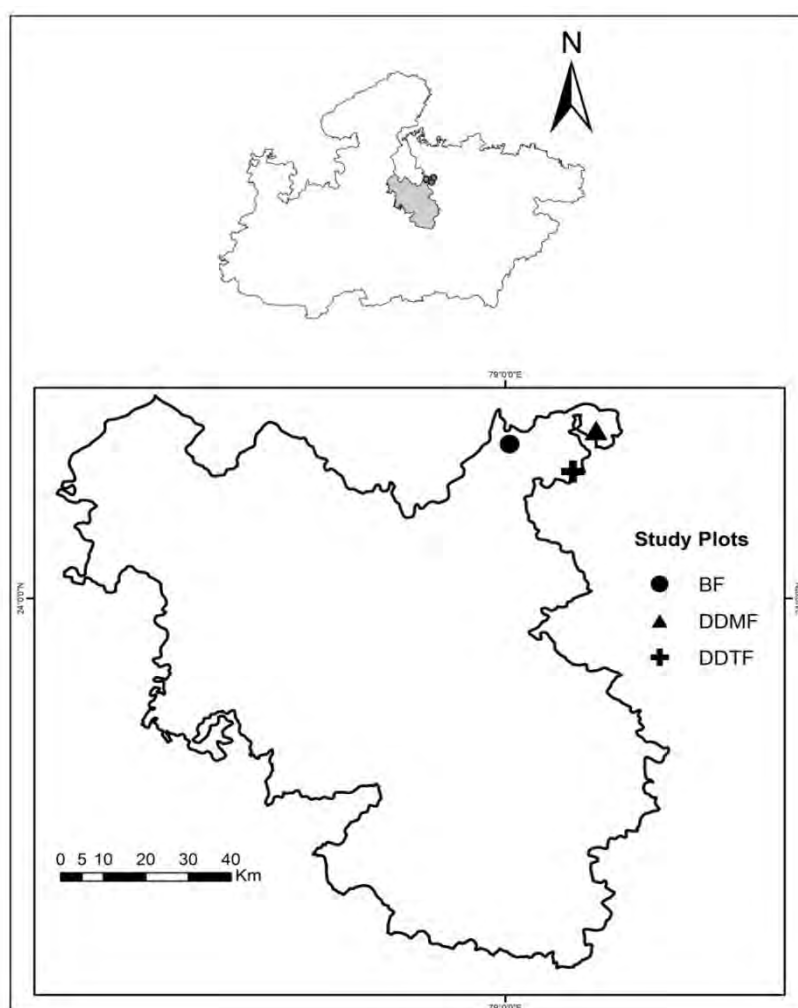


Fig. 1 Location of the study sites in Sagar, Madhya Pradesh.

2.3 Soil temperature (Ts) and soil moisture (Ms)

Soil temperature was measured using a digital soil thermometer at a depth of 0 - 10 cm adjacent to each soil CO₂ efflux beaker. Soil moisture was measured by gravimetric method at 0 to 10 cm depth adjacent to each CO₂ efflux beaker.

2.4 Statistical analyses

One-way ANOVA was used to test the differences between mean soil CO₂ efflux, soil temperature, and soil moisture among the three forest types, and Turkey's HSD test was applied whenever the ANOVA was significant. The level of significance for the analyses was set at $p < 0.05$ (Zhang, 2022). Linear correlation/regression analyses were used to examine the relationship of soil CO₂ efflux with soil temperature and soil moisture. SPSS 20.0 software was used for all statistical analyses.

3 Results

3.1 Monthly and seasonal variations in the soil CO₂ efflux

Soil CO₂ efflux varied significantly among all the forest types ($p < 0.001$). The monthly soil CO₂ efflux ranged from 116.3 to 626.0 mg CO₂/m²/h, with an average value of 338.6 ± 132 mg CO₂/m²/h. Soil CO₂ efflux varied significantly ($p < 0.001$) among the months. The mean monthly soil CO₂ efflux was measured to be 335.5, 385.4 and 294.7 mg CO₂/m²/h in site I, site II and site III respectively. The highest monthly soil CO₂ efflux was observed in July (626.0 mg CO₂/m²/h), compared to all other months in all the forest sites. Among the three study sites, site II showed a higher rate of soil CO₂ efflux compared to the other two study sites (Table 2).

Table 2 Mean soil respiration (mg CO₂/m²/h) under three different dry deciduous forest types in Sagar, Madhya Pradesh, India.

Months	Soil CO ₂ efflux (mg CO ₂ /m ² /h)			P-value
	DDTF	DDMF	BF	
January	142.8 ^a ± 6.9	162.8 ^b ± 10.9	122.3 ^c ± 5.0	0.001
February	238.4 ^a ± 10.8	267.1 ^b ± 8.5	190.9 ^c ± 4.1	0.001
March	250.1 ^a ± 4.5	277.1 ^b ± 16.8	226.4 ^c ± 8.5	0.001
April	335.7 ^a ± 9.4	370.9 ^b ± 15.9	270.5 ^c ± 3.2	0.001
May	337.1 ^a ± 6.9	339.5 ^b ± 13.8	281.7 ^b ± 3.0	0.001
June	418.5 ^a ± 7.9	467.1 ^b ± 7.4	373.4 ^c ± 11.3	0.001
July	563.5 ^a ± 4.8	614.1 ^b ± 9.1	489.5 ^c ± 9.7	0.001
August	529.8 ^a ± 5.2	603.5 ^b ± 6.7	478.3 ^c ± 32.7	0.001
September	431.9 ^a ± 20.6	530.6 ^b ± 18.3	371.9 ^c ± 9.1	0.001
October	337.1 ^a ± 7.0	456.3 ^a ± 25.2	317.6 ^b ± 7.7	0.001
November	231.1 ^a ± 9.8	303.3 ^a ± 9.3	220.5 ^b ± 9.0	0.001
December	210.1 ^a ± 7.6	232.9 ^b ± 3.2	193.6 ^c ± 7.1	0.001

Soil CO₂ efflux showed a strong seasonal pattern across all the forest sites peaking during monsoon season (mean 477 ± 57.3 mg CO₂/m²/h; range: 396.4 - 559.2 mg CO₂/m²/h), followed by summer (mean 329 ± 31.5 mg CO₂/m²/h; range: 286.2 - 377.1 mg CO₂/m²/h). The lowest values were observed during winter (mean 209.6 ± 24.9 mg CO₂/m²/h; range: 172.9 - 246.4 mg CO₂/m²/h, Table 3).

Table 3 Seasonal variation in soil respiration (RS), soil temperature (TS) and soil moisture (SM) under three different dry deciduous forest types in Sagar, Madhya Pradesh, India.

Forest	Season	Soil respiration (mg CO ₂ m ² /h)	P-value	Soil temperature (°C)	P-value	Soil moisture (%)	P-value
DDTF	Monsoon	465.6 ^b ± 88.5	0.001	29.4 ^b ± 3.2	0.001	16.6 ^b ± 2.0	0.001
	Winter	205.6 ^c ± 37.7		21.0 ^c ± 2.8		9.9 ^b ± 0.4	
	Summer	363.7 ^a ± 68.6		32.7 ^a ± 4.1		11.4 ^a ± 1.8	
DDMF	Monsoon	551.1 ^b ± 63.5	0.001	29.2 ^b ± 3.6	0.001	17.8 ^b ± 0.9	0.001
	Winter	241.5 ^c ± 51.8		20.0 ^b ± 3.0		10.9 ^c ± 0.4	
	Summer	288.0 ^a ± 53.4		32.2 ^a ± 4.4		9.6 ^a ± 2.1	
BF	Monsoon	414.4 ^b ± 72.3	0.001	27.5 ^a ± 3.3	0.001	16.1 ^a ± 1.6	0.001
	Winter	181.8 ^c ± 36.3		19.8 ^b ± 3.1		9.7 ^b ± 0.6	
	Summer	335.4 ^a ± 59.6		33.0 ^a ± 3.7		11.5 ^a ± 1.7	

Table 4 Mean soil moisture (%) under three different dry deciduous forest types in Sagar, Madhya Pradesh, India.

Months	Soil moisture (%)			P-value
	DDTF	DDMF	BF	
January	9.5 ^a ± 0.2	10.6 ^b ± 0.3	8.9 ^c ± 0.1	0.001
February	10.6 ^a ± 0.3	10.6 ^b ± 0.2	9.5 ^b ± 0.2	0.001
March	9.9 ^a ± 0.2	9.8 ^b ± 0.2	8.8 ^b ± 0.2	0.001
April	11.1 ^a ± 0.2	11.2 ^b ± 0.3	9.2 ^b ± 0.3	0.001
May	10.5 ^a ± 0.2	10.2 ^b ± 0.3	7.5 ^b ± 1.5	0.001
June	14.4 ^a ± 0.3	14.3 ^b ± 0.3	13.0 ^b ± 0.4	0.001
July	18.4 ^a ± 0.2	18.9 ^b ± 0.1	17.8 ^c ± 0.2	0.001
August	17.8 ^a ± 0.3	18.4 ^b ± 0.3	16.8 ^c ± 0.2	0.001
September	16.8 ^a ± 0.2	17.5 ^a ± 0.6	16.6 ^b ± 0.3	0.001
October	13.2 ^a ± 0.2	16.5 ^a ± 0.3	13.4 ^b ± 0.3	0.001
November	9.9 ^a ± 0.2	11.6 ^b ± 0.3	10.5 ^c ± 0.2	0.001
December	9.7 ^a ± 0.2	10.8 ^a ± 0.2	10.0 ^b ± 0.2	0.001

Mean value ± standard error (SE). Different letter (s) in the same row indicates significant differences at $p < 0.05$.

3.2 Soil temperature (T_S) and soil moisture content (M_S)

Mean monthly T_S and M_S showed significant differences among the forest sites ($p < 0.001$) (Tables 4 and 7.4). The average monthly T_S was 27.2°C. The maximum T_S was recorded in site I (37.2°C) during June, whereas, the minimum temperature was recorded under site II (15.3°C) during the month of January. The average monthly M_S was 12.6%. The highest monthly M_S was observed in site II (19.1%) in July, whereas the lowest M_S was measured in site III (6.1%) in May over the study period (Table 4).

Soil temperature showed a strong seasonal pattern across all the forest sites with the maximum during summer (25.7 - 37.2°C, Table 5) followed by monsoon and winter seasons, whereas soil moisture (%) showed a different trend. The highest M_S value was generally observed during monsoon when compared to other seasons (Table 5).

Table 5 Mean soil temperature (°C) under three different dry deciduous forest types in Sagar, Madhya Pradesh, India.

Forest type	Soil temperature (°C)			P-value
	DDTF	DDMF	BF	
January	17.9 ^a ± 0.5	15.8 ^b ± 0.3	16.6 ^c ± 13.3	0.001
February	25.2 ^a ± 0.2	24.0 ^a ± 0.5	24.5 ^b ± 0.3	0.001
March	27.9 ^a ± 0.5	26.1 ^a ± 0.2	26.4 ^b ± 0.2	0.001
April	31.0 ^a ± 0.4	32.8 ^b ± 0.6	29.3 ^c ± 0.3	0.001
May	36.1 ^a ± 0.4	35.3 ^b ± 0.3	36.4 ^b ± 0.3	0.001
June	37.0 ^a ± 0.1	36.7 ^a ± 0.3	36.5 ^b ± 0.3	0.003
July	34.1 ^a ± 0.4	33.9 ^b ± 0.4	32.4 ^b ± 0.3	0.001
August	30.6 ^a ± 0.7	30.7 ^b ± 0.8	28.2 ^b ± 0.2	0.001
September	27.0 ^a ± 0.4	27.9 ^b ± 0.5	25.5 ^c ± 0.4	0.001
October	26.0 ^a ± 0.9	24.2 ^a ± 0.2	23.7 ^b ± 0.4	0.001
November	21.7 ^a ± 0.4	20.9 ^a ± 0.5	20.4 ^b ± 0.2	0.001
December	19.1 ^a ± 0.5	19.4 ^b ± 0.2	17.5 ^b ± 0.3	0.001

Mean value ± standard error (SE). Different letter(s) in the same row indicates significant differences at $p < 0.05$.

3.3 Annual soil CO₂ efflux

The cumulative annual soil CO₂ efflux from April to December for each forest type was calculated as the sum of values of all the months and it ranged between 3536.8 and 4625.2 mg CO₂/m²/yr, with an overall mean value of 4062.7 mg CO₂/m²/yr across all the studied sites. The highest cumulative CO₂ efflux was observed in site II (4625.2 mg CO₂/m²/yr), followed by site I (4026.2 mg CO₂/m²/yr) and site III (3536.8 mg CO₂/m²/yr, Fig. 2).

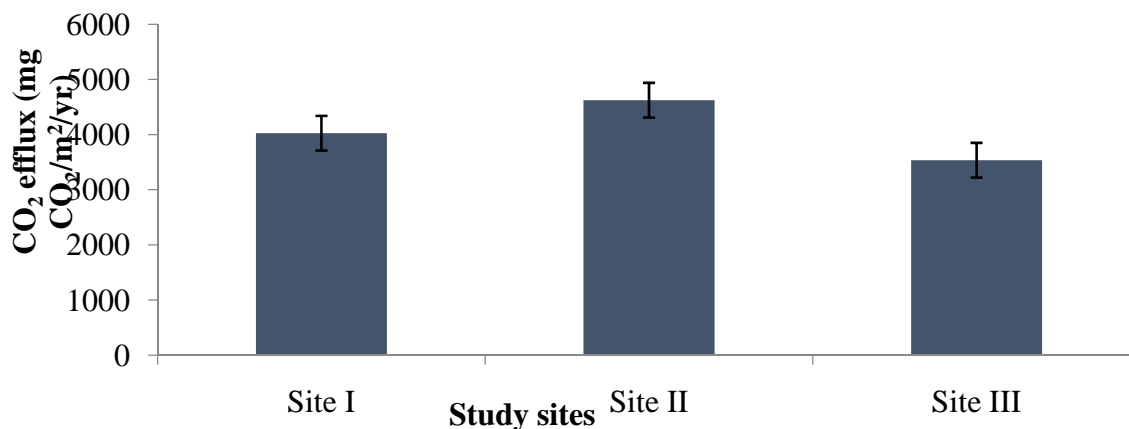


Fig. 2 Cumulative annual soil CO₂ efflux (mg CO₂/m²/yr).

3.4 Relationship between soil CO₂ efflux and environmental variables

Significant positive correlations were observed between soil CO₂ efflux and soil moisture in study site I ($R^2 = 0.88$, $p < 0.001$), study site II ($R^2 = 0.85$, $p < 0.001$) and study site III ($R^2 = 0.75$, $p < 0.001$, fig 3). Positive correlation was also observed between soil CO₂ efflux and soil temperature in study site I ($R^2 = 0.53$, $p < 0.001$), study site II ($R^2 = 0.44$, $p < 0.001$) and study site III ($R^2 = 0.39$, $p < 0.001$, Fig 4).

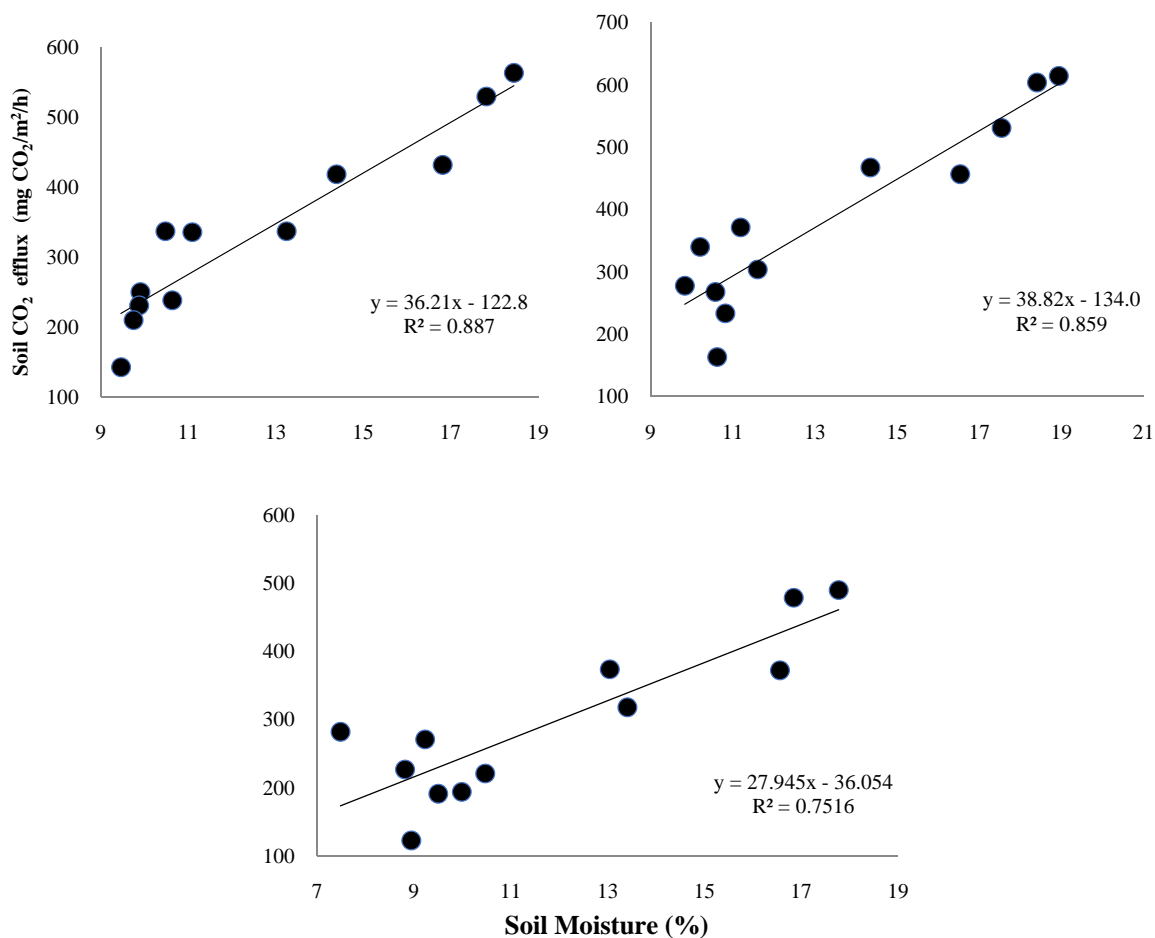


Fig. 3 Relationship between soil CO₂ efflux (month wise) and soil moisture in three studied sites.

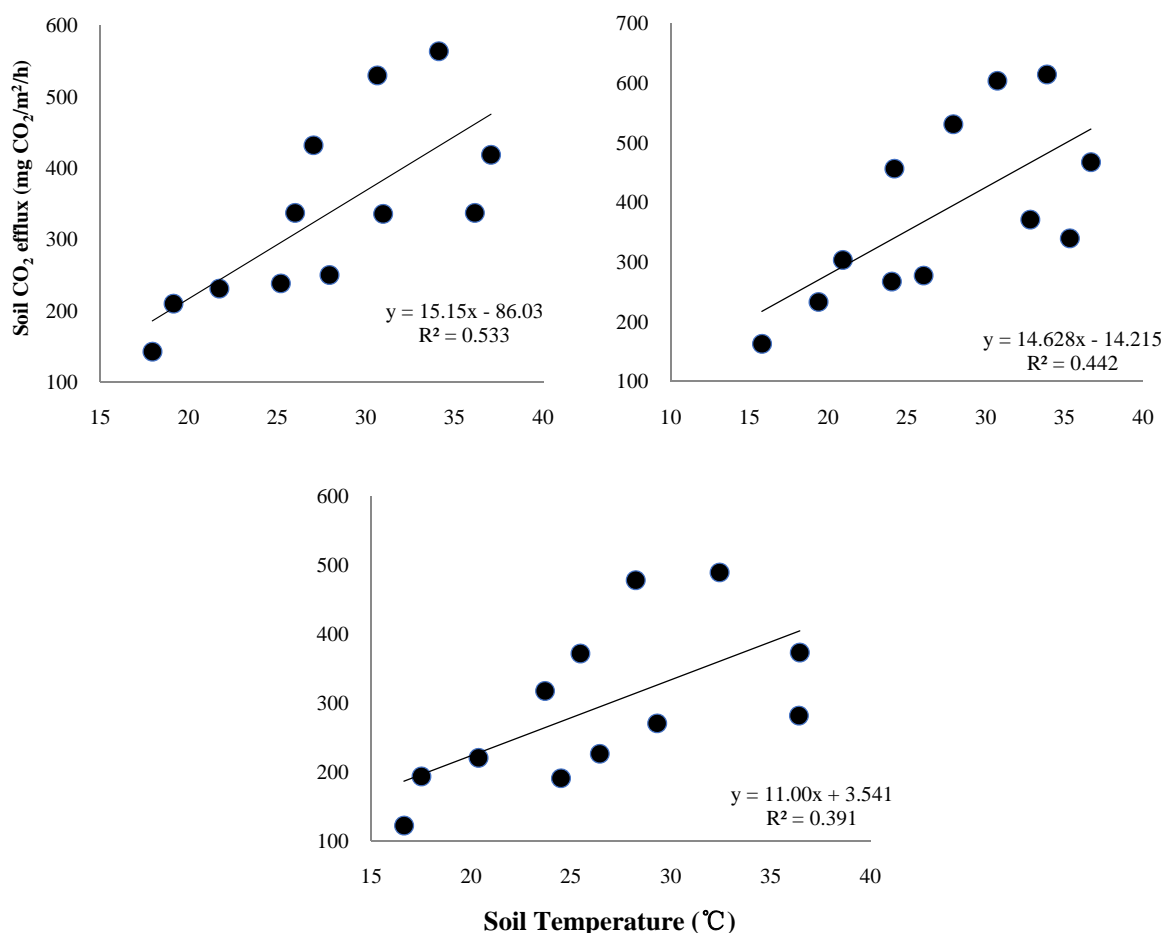


Fig. 4 Relationship between soil CO₂ efflux (month wise) and soil temperature in three studied sites.

4 Discussion

In this study, both monthly and seasonal variations in soil CO₂ efflux were observed in all the three tropical dry deciduous forest sites. The soil CO₂ efflux reached its peak during July and attained its minimum during January in all the study sites, which may be because of the similar climatic conditions in all the three sites. A strong seasonal variation in soil respiration has been observed across all the three study sites, with higher efflux rate during the monsoon season (July-October). The reason for the high rates of soil CO₂ efflux in the study sites during monsoon season could be the displacement of air rich in CO₂ from within the soil due to soil microbial activity which oxidizes the C dissolved in water. Soil contains small pores, which are usually filled with CO₂-rich air, when the soil is saturated; the small pores with air are substituted by water, producing a CO₂ efflux from the soil. The soil CO₂ efflux can therefore enhance rapidly during heavy rainfall (Lee et al., 2002; Luo et al., 2006). Increase in organic matter decomposition along with the proliferation in soil microbial (fungal, bacterial and actinomycetes) activity during the monsoon season which has frequent spells of rainfall with short periods of drying is regarded as an important reason for increased rates of soil respiration during rainy months (Jha and Mohapatra, 2011; Pandey et al., 2010). In all the three study sites, the minimum rate of soil CO₂ efflux during the winter season (November-February) may be attributed due to variations in the incoming solar radiation, temperature and day-length which result in low soil temperature and slow microbial activities (Dar et al., 2015; Han et al., 2012; Lloyd and Taylor, 1994; Pandey et al., 2010).

In this study, site II (DDMF) showed the highest rate of soil CO₂ efflux when compared with other two study sites. Higher soil CO₂ efflux in the study site II (DDMF) could be due to the higher tree density and floristic diversity. Floristic diversity and composition can influence the rate of soil CO₂ efflux by affecting soil microclimate and composition, the quantity and quality of detritus material, and the overall rate of root respiration (Raich and Tufekcioglu, 2000; Lee et al., 2010; Dias et al., 2010). Variations in soil CO₂ efflux due to differences in floristic diversity and vegetation attributes have been documented by several researchers (Dias et al., 2010; Han et al., 2012; Houghton et al., 2012; Raich and Tufekcioglu, 2000; Yan et al., 2011). Vegetation composition and stand structure are known to influence soil ecological properties such as soil temperature (Savva et al., 2010), soil moisture (Buytaert et al., 2006; Nosetto et al., 2005; Wang et al., 2012), and quality and quantity of soil organic matter (Smith et al., 2014).

In general, the upper soil layer is the zone with maximum root concentration and soil biological activity (Tripathi and Singh, 1992; Tripathi et al., 1999), so the results obtained in the present study were found to be fairly comparable with other studies. The range of soil CO₂ efflux recorded in the present study (116.3 to 626.0 mg CO₂/m²/h) is similar with other reports from different forests Devi and Singh, 2016 (195.71 - 345.98 mg CO₂/m²/h); Devi and Yadava, 2009 (138.49–250.94 mg CO₂/m²/h); Fernandes et al., 2002 (216 - 510 mg CO₂/m²/h); Gupta and Singh, 1981 (44 - 448 mg CO₂/m²/h); Joshi et al., 1991 (62.4 - 362 and 37.6 - 282 mg CO₂/m²/h); Laishram et al., 2002 (368–634.23 mg CO₂/m²/h); Lamotte, 1975 (330 mg CO₂/m²/h); Leith and Quелlette, 1962 (251 mg CO₂/m²/h); Medina and Zelwar, 1972 (234 - 511 mg CO₂/m²/h); Monteith et al., 1964 (312 - 500 mg CO₂/m²/h); Pandey et al., 2010 (102 - 320 and 99 - 543 mg CO₂/m²/h); Redmann, 1978 (0 - 433 mg CO₂/m²/h); Schwendenmann et al., 2003 (430–675 mg CO₂/m²/h); Sundarapandian and Dar, 2013 (126 - 427 and 182 - 646 mg CO₂/m²/h); Tewari et al., 1982 (101.3 - 270 mg CO₂/m²/h); Thokchom and Yadava, 2014 (124.33 - 586.03 mg CO₂/m²/h).

Climatic factors, in particular, soil temperature and moisture are known to be the principal governing factors of soil CO₂ efflux (Rodtassana et al., 2021; Sundarapandian and Kirthiga, 2011). In the present study, soil temperature showed a significant positive correlation with soil CO₂ efflux in all the three study sites. A similar trend has been reported by several researchers (Cao et al., 2004; Fang and Moncrieff, 2001; Jha and Mohapatra, 2011; Peri et al., 2015; Rubio and Detto, 2017). Soil temperature influences soil CO₂ efflux rates by increasing the metabolic activity of soil microbes upto an optimal threshold, after which the soil microbial activity and rate of soil CO₂ efflux begin to decline (Conant et al., 2004; Meena et al., 2020). Higher soil temperatures lead to more rapid decay of soil organic matter, lesser C accumulation in the slow and passive pools, and a larger loss of C via respiration, provided there is adequate moisture (Okello et al., 2023; Canadell et al., 2007). Hence, soil temperature is often considered as a major factor that drives seasonal variations in soil CO₂ efflux in forest ecosystems (Rodtassana et al., 2021, Arroyo and Wood, 2021; Reichstein et al., 2003; Bahn et al., 2008).

Besides soil temperature, soil moisture is another main factor that greatly regulates soil CO₂ efflux (Akburak and Makineci, 2013; Dar et al., 2015). In the present study, soil CO₂ efflux was found to be significantly positively correlated with soil moisture in all the three study sites. Similar results were obtained in other studies as well (A'Bear et al., 2014; Dar et al., 2015; Devi and Yadav, 2008; Jangra et al., 2011; Londo et al., 1999; Morén and Lindroth, 2000; Ohashi et al., 1999; Soe and Buchmann 2005; Steinweg et al., 2013; Sundarapandian and Dar, 2013). In this study, soil CO₂ efflux was found to be highest during the monsoon season, when soil moisture was also at its peak. Soil moisture acts as a medium for solubilization of the organic substrates and microbial activity (Koizumi et al., 1999). Hence, soil moisture influences different processes like soil organic matter decomposition, soil C mineralization, microbial activity, etc. which in turn, influence soil respiration (Bao et al., 2016; Meena et al., 2020; Zhang et al., 2019). Although many studies

indicated a strong correlation between soil CO₂ efflux and soil organic carbon (Dube et al., 2009; Iqbal et al., 2010; Wang et al., 2013), a non-significant, yet positive relationship between the two variables was observed in this study. This is so because soil CO₂ efflux is mainly determined by living plant roots, and organic matter mineralization by soil microbial activity (Buchmann, 2000).

5 Conclusions

The present study revealed the monthly and seasonal variations in soil CO₂ efflux in the three tropical dry deciduous forest types in Madhya Pradesh. It is also evident that the soil CO₂ effluxes in the three study sites are mainly controlled by soil temperature and soil moisture. Changes in temperature and precipitation regimes as a result of climate change could affect the rates of soil CO₂ efflux in tropical dry forests and alter the global C balance.

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